

Comparison of Some Mathematical Models to Calculate Evapotranspiration in Contrasting Regions of Iraq

Dher I. Bakr¹, Jasim Al-Khalidi¹ and Azhar S. Hadi²

¹Department of Physics, College of Sciences, University of Diyala, Diyala, Iraq. ²Department of Geography, College of Education for Human Sciences, University of Diyala, Diyala, Iraq.

*Corresponding Author: dharentisar78@gmail.com Received: June 30, 2020; Revised: August 16, 2020; Accepted: December 23, 2020

Abstract

Evaporation is one of the major components of hydrologic cycle and water balance. The data required for its accurate estimation are commonly available only at widely spaced weather stations. Evapotranspiration can be estimated using mathematical models based on atmospheric variable, which can be measured directly. However, these mathematical models differ among themselves in terms of accuracy of the results due to the different determinants of each mathematical model. In addition, the accuracy of the model varies from region to region. The current study aims to compare some mathematical models performance in selected regions in Iraq. The results showed that Penman-Monteith model was the best model for calculating Evapotranspiration in Iraq as it achieved values close to the observed values of evapoetranspiration with a high correlation and an absolute error coefficient ratio is small. In comparison, Ivanov model was the least accurate model with a smaller correlation values and higher absolute error rate values than the other methods . The data used in this study are a time series of the evapotranspiration values of agricultural stations. These data are measured by means of lysrometers. The model that uses the largest number of atmospheric variables that influence on Evapotranspiration calculation is the best.

Keywords: Evapotranspiration; Penman-Monteith; Thornthwaite; Khrrufa; Ivanov; Correlation; Mean Absolute Error.

1. Introduction

Water resources management is a critical demand for increased agricultural production in arid regions where food insecurity has become a major concern (De Zeeuw *et al.*, 2009). The water cycle is the main process in the different iteractions and for different scales of the earth system. However, Earth surface evapotranspiration (ET), as an important part of the water cycle, is the main method of water consumption in the Earth system, as it plays an important role in the global and regional climate of the lands (Wang *et al.*, 2012; Jung *et al.*, 2010). Farmland ET refers to the total water exchanged with the air by plants and the Earth's surface respectively.

ET around the world consumes about 60% of the rain and 99% of the water in the farmland system (Kite, 2000). With global climate change, evapotranspiration, is a prerequisite for understanding the interactions between soil, plants and the atmosphere (Nemani *et al.*, 2002). ET consists of evaporation and transpiration. The first is the process by which water transfers from the surface, liquid to vapor, and enters the atmosphere, and the second is the process by which water in plants passes through the stomata and spreads outward (Lulu *et al.*, 2012). As for reference evapotranspiration (ETo) represents the maximum amount of steam

that can be released from a given area of the earth to the atmosphere is under the influence of the weather factors specific to that region (Ammar and Haidar, 2013). Evaporation is an component not only in the surface heat balance but also in the hydrological cycle (Jingxin *et al.*, 2013). It has the functions of regulating temperature, increasing humidity, affecting the environmental ambience, and thus affects the sustainable development of society and economy (Liu *et al.*, 2013). Farmland irrigation and ET can use and insert 60% and 80% of the net radiation during the growing season, receptively (Suyker *et al.*, 2008).

ET is a function of a number of environmental variables including temperature, precipitation, radiation, wind velocity and humidity (Lulu *et al.*, 2012; Liu *et al.*, 2013). Therefore, varying ET good indicator of climate change drivers (Chattopadhyay *et al.*1997).

In the literature, different works have used a variety of methods for estimating ET (Roderick and Farquhar, 2002). Lysimeters have been applied to determine ET of mass balance for shallow groundwater environments. However, it has been found that this method might give false readings due to air entrapment (Fayer and Hillel, 1986), as well as fluctuating by the water table (Yang *et al.*, 2000). In addition, some methodologies such as the energy budget and eddy correlation are also restricted to calculating net ET(Net evapotranspiration). Net ET occurs when the rate of evaporation exceeds the rate of condensation (Summer, 2006).

The Penman–Monteith method was found better than the Kharrufa, Samani, Hargreaves models and modified Blaney–Criddle methods to estimate the ET in Mosul city in Iraq (Al-Rijabo *et al.*, 2008). In Karbala city (Aljumaili *et al.*, 2014), the ET was estimated using the Food and Agriculture Organization (FAO) Penman - Monteith, Penman -Kimberly, Jensen - Haise and Hargreaves models; the first of those methods was more consistent with observed data than the others. Studies of Al - Sudani (2018) and Jawad (2016) adopted the Thornthwaite equation to estimate the ET in Iraq. The results were dependable for calculating ET.

In this study, monthly time series of observation and estimated data are compiled and tested by the Pearson Index (R^2) and the Mean Absolute Error (MAE). We used four methods over the four agriculture stations to investigate the best one to characterize the ET in Iraq.

2. Methodology

2.1 Study area

The monthly data of the climatic parameters from the Iraqi Meteorological Organization are used in to estimate the values of ET by the four methods (FAO Penman - Monteith, Thornthwaite, Khrrufa, and Ivanov), also the monthly real values of ET are provided through the same source. The chosen study area is represented by four meteorological stations distributed over four Iraqi governorates; Mosul, Baghdad, Anbar and Basrah as shown in (Figure 1), and located in different regions in terms of longitude and latitude, as shown in (Table 1). These governorates have the advantage that they differ in terms of geographical characteristics. In addition, the variation in the values of atmospheric elements that affect the value of evaporation, as well as these governorates have the advantage that they include all climatic patterns for all governorates of Iraq.

 Table 1. The geographical and weather characterization of some climatic parameter for the study stations

Stations	Altitude	Longtude	Longtude	Avg. temperature	Wind	Monthly total	Relative
	(m)	(°E)	(°N)	(C)	velocity	solar rad.	humidity
					(m/sec)	(Mj/m^2)	(%)
Baghdad	31	44.23	33.32	23.37	1.62	18.6	42.8
Mosul	223	43.16	36.33	21.95	1.2	17.3	48.98
Al-Anbar	30	43.97	32.55	23.38	1.67	18.68	38.6
Basrah	2	47.45	30.93	13.13	3.78	25.87	38.3



Figure 1. Iraq map, Baghdad, Mosul, Anbar and Basrah weather stations locations.

2.2 Methods

In this study, four mathematical models FAO Penman - Monteith, Thornthwaite, Khrrufa, and Ivanov methods are used to estimate the values of the Evapotranspiration. Where data for the region is entered into the four mathematical models followed by a comparison of the evaporation values of the Evapotranspiration that were estimated for each model with values Evapotranspiration observation. The data of the observed values used in this study were obtained from the Agricultural Meteorological Network at the Iraqi Ministry of Agriculture. Regional data were analyzed using Pearson correlation statistics with error checking the mean absolute error (MAE). This process was applied to all four chosen agricultural areas.

Because the calculation of Evapotranspiration using Evapotranspiration measurement devices is not accurate in addition to neglecting transpiration from the plant (Ammar and Haidar, 2013), so Khrrufa, Thornthwaite, Hargreaves and Ivaov have found mathematical equations by which the transpiration evaporation can be estimated, and many equations have appeared in the calculation of transpiration evaporation on temperature, soil heat, wind velocity, vapor pressure-and sun shine duration, some of which relied on temperature and humidity like (Ivanov) and others depend on the temperatures and rates of solar brightness (Kinjo, 2009), such as the modified lamb. These equations were adopted in estimating the evaporation values of latent transpiration in Iraq due to their ease of application and the availability of climatic data required in the model.

2.2.1 FAO Penman-Monteith method Its mathematical formula (Biesdore et al., 2017)

$$ET_o = \frac{\left[0.408\nabla(Rn - G) + r\left\{\left(\frac{900}{T273}\right)U2(es - ea)\right\}\right]}{\Delta + r(1 + 0.34U2)}$$
(1)

Where:

ET₀: reference evapotranspiration (mm day⁻¹). G: soil heat flux (MJ/m² day). Rn: net radiation (MJ/m² day)

 Δ : slope of the saturation vapor pressure-temperature curve (Pa/ °C).

r : humidity constant.

T: temperature (C)

U2: wind velocity on hight 2meter (m/sec)

 e_s :saturated water vapor pressure (kPa).

 e_a : Actual water vapor pressure (kPa).

2.2.2 Thornthwaite method (Saud et al., 2014)

$$ET_o = 16 * \left(10\frac{T}{I}\right)^a * \mu \frac{N}{360}$$
 (2)

Where:

E: is monthly potential evapotranspiration (mm/month)

T: monthly mean temperature (C)

I: the empirical annual heat index, the sum of 12 monthly index values i. The value of i for each month is derived from mean monthly temperatures according to the formula:

 $ij = 0.09^{*}(Tj)1.5$, where subscript j indicates the specific month under investigation,

 $\boldsymbol{\mu} {:}$ the number of days in the month

N: the mean number of daylight hours in a particular month

a: an empirically derived exponent which is a function of I, and is given by the formula: a = 0.016*I + 0.5.

2.2.3 Khrrufa Equation (Chen et al., 2005)

$$ET_o = 0.34 p T^{1.3} (3)$$

T is mean temperature (C), p is sunrise hours (sun shine duration)

2.2.4 Ivanov method (Cunha et al., 2017)

 $ET_o = 0.0018 * (25 - T) * (100 - RH)$ (4)

T is mean monthly temperature (C), RH is Relative Humidity

For the purpose of knowing the efficacy of the mathmatical models and evaluate the performance of these models in ET Calculation. The comparison is made between the estimated and measured evapotranspiration using the Class A pan method values. Class A evaporation ponds are widely used for estimating the ETo reference evaporationtranspiration of green grass reference plant (Snyder, 1992)

Several performance criteria were used including Pearson type coefficient of determination index (R²) and Mean Absolute Error (MAE), these criteria are defined as: (a) Pearson type coefficient of determination index, R^2 (Mondal and Mondal, 2017)

$$R^{2} = 1 - \frac{\sum_{i=1}^{n} (y_{i} - \hat{y})^{2}}{\sum_{i=1}^{n} (y_{i} - \bar{y})^{2}}$$
(5)

Where: R^2 is Pearson type coefficient of determination index

yi : observation values, \bar{y} : average of values, n: number of values, \hat{y} : estimated values of (ET) The value of (ET) represents the estimated values of evapotranspiration, which can be estimated using the mathematical models that were adopted in this study.

(b) Mean Absolute Error (MAE): measures the average magnitude of the errors in a set of predictions, without considering their direction. It's the average over the test sample of the absolute differences between prediction and actual observation. and its mathematical formula (Wang and Lu, 2018)

$$MAE = \frac{1}{n} \sum_{i=1}^{n} \left| \frac{\hat{y} - y_i}{y_i} \right|$$
(6)

 y_i and \hat{y} :, are the observation and estimated values respectively

Mean Absolute Error (MAE) represents the mean deviation between two values, its optimal value is zero. These statistics are between the estimated values and the observed real values. It is worth noting that each M.A.E measures errors as it gives an indication of the difference between the real values and the estimated values. This has been used to test the accuracy of many mathematical models.

The MAE, R² values are important criteria in assessing the efficiency of the method used to estimate evapotranspiration values.

3. Results

The evaluations of the four methods are emphasized through the comparison between the observed and estimated evapotranspiration time series. Pearson type coefficient of determination index (\mathbb{R}^2) and Mean Absolute Error (MAE) tests are used to determine the best model also.

3.1 Comparison between observation and estimation evapotranspiration data for Baghdad.

Time series were used for the observed and estimated evapotranspiration data to find out the nature of the similarities or differences between them (Fig. 2a-c). The time series estimated by Penman-Monteithand and Thornthwaite methods (Fig 2a,b) have better consequence with the observed data than the other two methods (Fig. 2c,d). The Pearson correlation index and Mean Absolute Error between the observed data and estimated values from the four methods are emphasized. For Baghdad station, the estimated values by Penman-Monteith (Table 2) method correlated strongly (0.92) with the observation data, the value of R² is significant and the MAE has very small value (0.20), these values refer to confirm with the similarity of the values calculation and observation time series. The Thornthwaite method is also significant for estimate the evapotranspiration for Baghdad station but less than the Penman-Monteithmethod, the MAE is (0.43). The R^2 value is (0.88) less than the other three method but the Thornthwaite method has less error than Khrrufa and Ivanov (Table 2). Khrrufa and Ivanov method have higher values of MAE than Penman-Monteithand and Thornthwaite that that mean these two methods not good to estimate the evapotranspiration in in Baghdad.



Figure 2. Observation and estimation of evapotranspiration in Baghdad station by (a) Penman, (b)Thornthwaite, (c) Khrrufa and (d) Ivanov methods

Table 2. Pearson index (\mathbb{R}^2) and Mean Absolute Error (MAE) for the Penman, Thornthwaite ,Khrrufa and Ivanov methods in Baghdad

Station	Method	\mathbb{R}^2	MAE
	FAO Penman-Monteith	0.92	0.20
Baghdad	Thornthwaite	0.88	0.43
	Khrrufa	0.91	0.84
	Ivanov	0.87	1.09

3.2 Comparison between observation and estimation evapotranspiration data for Mosul.

To evaluate the evapotranspiration values which estimated from the four methods and compare it with the observed values, the time series for estimated and observed data are configured (Fig. 3a-c), Pearson index and Mean Absolute Error are used also. Time series computed through Penman-Monteithand and Thornthwaite methods (Fig. 3a,b) have similar behavior more than the other two methods to estimate the evapotranspiration (Fig.3c,d). For Mosul station, the estimated values by Penman-Monteith (Table 3) method correlated strongly (0.99) with the observation data, the value of R^2 is significant and the MAE has very small value (0.26), these values are confirm with the similarity of two time series. The Thornthwaite method is also significant for estimate the evapotranspiration for Mosul station but less than the Penman-Monteith method, the MAE is (0.43). The R² value is (0.93) less than the other three methods but the Thornthwaite method has less error than Khrrufa and Ivanov (Table 3).Khrrufa and Ivanov method have higher values of MAE than Penman - Monteithand and Thornthwaite that reversed the efficiency of the Penman - Monteithand and Thornthwaite methods to represent the evapotranspiration in Mosul.



Figure 3. observation and estimation of evapotranspiration in Mosul station by (a) Penman, (b)Thornthwaite, (c) Khrrufa and (d) Ivanov methods

Table 3. Pearson index (R²) and Mean Absolute Error (MAE) for the Penman, Thornthwaite, Khrrufa and Ivanov in Mosul

Station	Method	\mathbb{R}^2	MAE
Mosul	FAO Penman-Monteith	0.99	0.26
	Thornthwaite	0.93	0.43
	Khrrufa		0.83
	Ivanov	0.94	0.80

3.3 Comparison between observation and estimation evapotranspiration data for Al-Anbar.

Time series were used for the observed and estimated evapotranspiration data to find out the nature of the similarities or differences between them (Fig. 4a-c)) . The time series estimated by Penman-Monteithand and Thornthwaite methods (Fig. 4a,b) have better representation for evapotranspiration than Khrrufa and Ivanov methods (Fig.4c,d). The Pearson index and Mean Absolute Error between the observed data and estimated values from the four methods are used to test the best method that could present the real data evapotranspiration. For Al-Anbar station, the estimated values by Penman-Monteith(Table 4) method correlated strongly (0.98) with the observation data, the value of R² is good and the MAE has very small value (0.25), these values confirm the similarity between the calculated and observed time series. The Thornthwaite method is also good for estimate the evapotranspiration for Al-Anbar station but less than the Penman-Monteith method, the MAE is (0.40). The \mathbb{R}^2 value is (0.94) less than the other three methods but the Thornthwaite method has less error than Khrrufa and Ivanov(Table 4). Khrrufa and Ivanov method have higher values of MAE than Penman-Monteith and Thornthwaite. Therefore, Penman-Monteith and Thornthwaite methods are the best in estimating the values of evapotranspiration in Al-Anbar.



Figure 4. observation and estimation of evapotranspiration in Al-Anbar station by (a) Penman, (b)Thornthwaite, (c) Khrrufa and (d) Ivanov methods

Table 4. Pearson index (R2) and Mean Absolute Error (MAE) for the Penman, Thornthwaite,Khrrufa and Ivanov methods in Al-Anbar

Station	Method	R ²	MAE
Al-Anbar	FAO Penman-Monteith	0.98	0.25
7 H 7 HIOU	Thornthwaite	0.94	0.40
	Khrrufa	0.97	0.83
	Ivanov	0.95	1.11

3.4 Comparison between observation and estimation evapotranspiration data for Basrah

Time series were used for the observed and estimated evapotranspiration data to find out the nature of the similarities or differences between them (Fig. 5a-c). The time series estimated by Penman-Monteith and Thornthwaite methods (Fig.5a,b) have better consequence with the observed data than the other two methods (Fig.5c,d). The Pearson index and Mean Absolute Error between the observed data and estimated values from the four methods are used to test the best method that could present the real data evapotranspiration. For Basrah station, the estimated values by Penman-Monteith (Table 5) method correlated strongly (0.94) with the observation data, the value of R^2 is good, the MAE has very small value (0.21). This indicates the convergence of the estimated with the observed values. The Thornthwaite method is also good for estimate the evapotranspiration but less than the Penman-Monteith method, the MAE is (0.47). The R² value is (0.95)(Table 4). Khrrufa and Ivanov methods have higher values of MAE than Penman-Monteithand and Thornthwaite that mean these two methods not good to estimate the evapotranspiration in Basrah station.



Figure 5. observation and estimation of evapotranspiration in Basrah station by (a) Penman, (b)Thornthwaite, (c) Khrrufa and (d) Ivanov methods

Table 5. Pearson index (R^2) and Mean Absolute Error (MAE)for the Penman, Thornthwaite,Khrrufa and Ivanov methods in Basrah

Station	Method	\mathbb{R}^2	MAE
	FAO Penman-Monteith	0.94	0.21
Basrah	Thornthwaite	0.95	0.47
	Khrrufa	0.96	0.85
	Ivanov	0.87	0.65

4. Conclusion

This study has contributed to evaluate the best method can use to estimate the evapotranspiration in Iraq. The main results of this study are presented as follows:

(1) The best method to estimate the evapotranspiration in Iraq is Penman-Monteith achieved values close to the observed values with high correlation and less ratio of the absolute error than the other models. In addition, the model related to the largest numbers of variables that influence the ET, either directly or indirectly.

(2) The Ivanov model is the least accurate model for the results because it achieved results with little correlation with the high absolute error ratio compared to other models in the region in which it was applied.

(3) The values of MAE test is more significant than R^2 in evaluating the evapotranspiration through the mathematical methods.

By relying on the four mathematical models and through data for the elements included in these models for the selected agricultural regions. It was found that there is a correlation between the real evapotranspiration values and the estimated values using mathematical models as shown in the figures 2 to 5.

The study noted that the values of these indicators are not equal, and this is due to the variation in climatic conditions and the geological characteristics of these areas, which affect the accuracy of the models

As these figures show that the relationship between the estimated values and the real values is not the same in relation to the mathematical model of the four chosen areas when calculating the evaporation values, the reference is succeeded, because mathematical models in general have determinants as they achieve acceptable values when applied in regions and achieve less accurate results in regions others in some cases, their results are not realistic in other regions due to the limitations of each model. The mathematical model that achieves acceptable results in one region does not have to achieve the same precision if it is applied in other regions. There is also a variation in accuracy for different mathematical models, if they are applied in the same area. That is, for each region it can be a mathematical model that achieves acceptable results more than the rest of the models where its determinants fit with the conditions of that region and it can be adopted in estimating the value of evaporation, we refer the reference to that region. But if the mathematical model achieves acceptable values in more than one region, then this means that the model is more comprehensive and its determinants fit all conditions of those areas.

From the values in Table 2 to 5, it appears that the FAO Penman-Monteith method is the best method to estimate the ET for all the four stations, this results is similar the finding of (Al-Rijabo et al., 2008; Aljumaili et al., 2014). The Thornthwaite model is a good method to estimate the ET in Iraq, this results is similar to the study of (Al-Sudani, 2018; Jawad, 2016). The main finding of this study is that the best model can be used to estimate the ET is the method that have the largest number of essential variables that influence the ET, directly or indirectly, as these elements act as indicators of estimated values, which is FAO Penman-Monteith model in our study.

Acknowledgments

The authors are grateful to the Iraqi meteorological organization for providing the data.

Conflicts of Interest

The authors declare no conflict of interest.

Reference

- Aljumaili KK., Al-Khafaji MS, Al-Awadi AT . Assessment of evapotranspiration estimation models for irrigation projects in Karbala, Iraq. Engineering and Technology Journal 2014; 32(5): 1149-1157.
- Al-Rijabo W, Jasim K., Shalal M. Study of reference evapotranspiration in Ninava Governorate using different mathematical models. AlRafedain Sciences Journal 2008; 19(3A): 154-173 (in Arabic).
- Al-Sudani H, Ilaibi Z. Study of morphometric properties and water balance using thornthwaite method in Khanaqin Basin, East of Iraq. Journal of University of Babylon for Engineering Sciences 2018; 26(3): 165-175.
- Ammar G, Haidar B . estimation of monthly reference evapotranspiration in Safita area by using artificial neural network. Tishreen University Journal for Research and Scientific Studies - Engineering Sciences Series 2013; 35(4): 28- 43.
- Biesdorf EM, Maciel EB, Teixeira MF, Teixeira JS, Salla PH, Pimertel LD, Imbuzeiro HA. Comparison of reference evapotranspiration estimates obtained by different methods in relation to the Penman-Montheith method in Cerrado Environment, Brazil. Journal of Experimental Agriculture International 2017; 17(1): 1-9.
- Chattopadhyay N, Hulme M. Evaporation and potential evapotranspiration in India under conditions of recent and future climate change. Agricultural and Forest Meteorology 1997; 87(1): 55-73.
- Chen JF, Yeh HF, Lee CH, Lo WC. Optimal comparison of empirical equations for estimating potential evapotranspiration in Taiwan. In: XXXI IAHR Congress 2005: 3867-3697.
- Cunha FF, Magalhaes FF, De Castro MA, De Souza EJ. Performance of estimative models for daily reference evapotranspiration in the city of Cassilândia, Brazil. EngenhariaAgrícola 2017; 37(1): 173-184.

- De Zeeuw H, Dubbeling M. Cities food and agriculture: Challenges and the way forward. RUAF Foundation, Leusden. 2009.
- Fayer MJ, Hillel D . Air encapsulation: Measurement in a field soil. Soil Science Society of America Journal 1986; 50(3): 568-572.
- Jawad LA. The use of geographic information system facilities to estimate the evapotranspiration in Iraq according to thornthwaite adjusted formula. Iraqi Journal of Science 2016; 57(4): 2566-2574.
- Jingxin LI, Wang SLI, Y SK, Xiang D . The evaporation variation and its Influence factors in Xi'ning of Qinghai province. Journal of Arid Meteorology 2013; 31(3): 497-504.
- Jung M, Philippe C, Seneviratnem SI, Sheffield J, Goulden ML . Recent decline in the global land evapotranspiration trend due to limited moisture supply. Nature 2010; 467(7318): 951-954.
- Kinjo A. Comparaison de quelques méthodes pour estimer l'evapotranspivation potentielle (ETP) dans la region de Lattaquie. Tishreen University Journal for Research and Scientific Studies - Biological Sciences Series 2010 ;3 2(1):10 -17. (in arabic)
- Kite G. Using a basin-scale hydrological model to estimate crop transpiration and soil evaporation. Journal of Hydrology 2000; 229(1): 59-69.
- Liu Z, Kennethg H, Xiaomao L, Xiaoguang Y. Negative effects of climate warming on maize yield are reversed by the changing of sowing date and cultivar selection in Northeast China. Global change biology 2013; 19(11): 3481-3492.
- Lulu S, Yunhe Y, Shaohong Wu. Advancements of the metrics of evapotranspiration. Progress in Geography 2012; 31(9): 1186-1195.
- Mondal S, Mondal H.Value for (2) in Statistical Analysis by Pearson Correlation Coefficient. Journal of Clinical and Dignostic Research 2017; 11(11): 1-2.

- Nemani R, Michael W, Peter T, Kenlo N, Swarna R, Jennifer J, Steven R. Recent trends in hydrologic balance have enhanced the terrestrial carbon sink in the United States. Geophysical Research Letters 2002; 29(10): 1061-1064.
- Roderick ML, Farquhar GD. The cause of decreased pan evaporation over the past 50 years. American Association for the Advancement of Science 2002; 298(5597): 1410-1411.
- Saud A, Kamel A, Azlin M, Abdullah R. Temporal and spatial variability of potential evapotranspiration in semi-Arid Region: Case study the Valleys of Western Region of Iraq. International Journal of Engineering Science and Technology 2014; 6(9): 653.
- Snyder, R.L. Equation for evaporation pan to evapotranspiration conversion. Journal of Irrigation and Drainage Engineering of ASCE 1992; 118(6): 977-980.

- Summer, David M. Adequacy of selected evapotranspiration approximation for hydrologic simulation. Journal of the American Water Resources Association 2006; 42(3): 699-711.
- Suyker AE, Verma SB. Interannual water vapor and energy exchange in an irrigated maizebased agroecosystem. Agricultural and Forest Meteorology 2008; 148(3): 417-427.
- Wang K, Dickinson RE. A review of global terrestrial evapotranspiration: Observation, modeling, climatology and climatic variability. Reviews of Geophysics 2012; 12(373):1-54.
- Wang W, Lu Y. Analysis of the mean absolute error (MAE) and the root mean square error (RMSE) in assessing rounding model. In: IOP Conference Series: Materials Science and Engineering 2018; 324(1):1-10.
- Yang J, Baoqing L, Liu S. A large weighing lysimeter for evapotranspiration and soil-water-groundwater exchange studies. Hydrological processes 2000; 14(10):1887-1897.